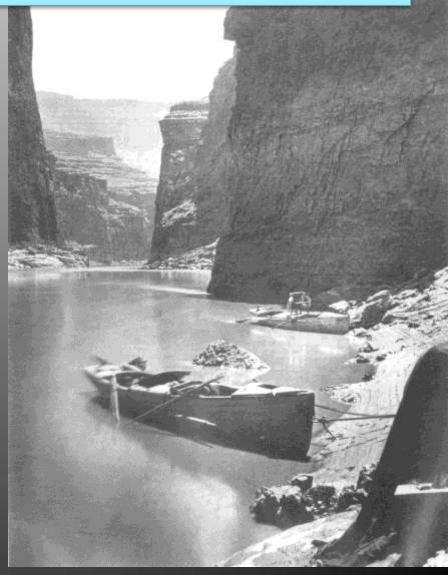
A Conversation about Rivers and Dams ...

John (Jack) C. Schmidt Grand Canyon Monitoring and Research Center jcschmidt@usgs.gov



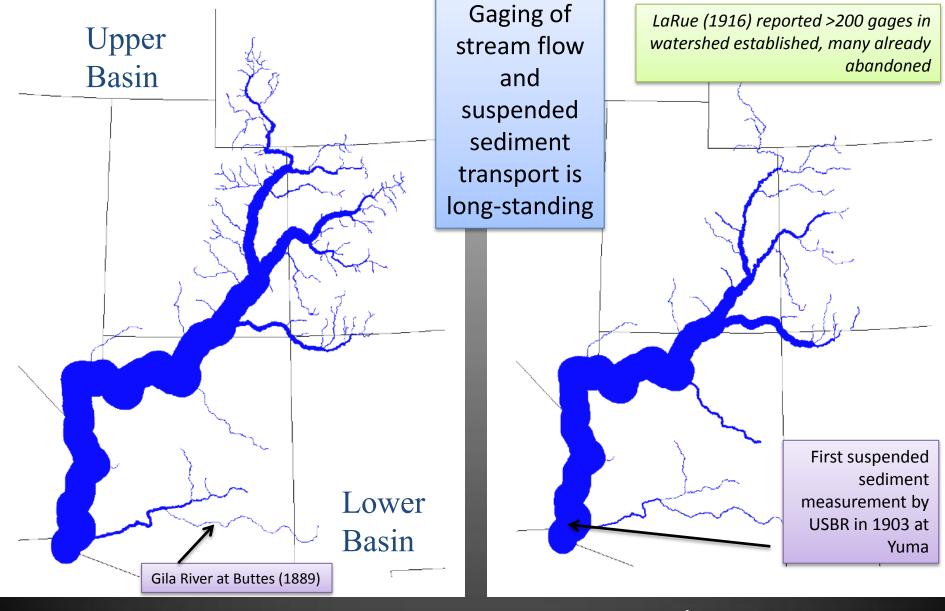
First scientific observations in Grand Canyon made by Powell (1875, 1895). Early photographs are an essential benchmark for environmental monitoring.





Badger Creek Rapids, 1920s

Marble Canyon, 1872



water

the upper 15% of the basin

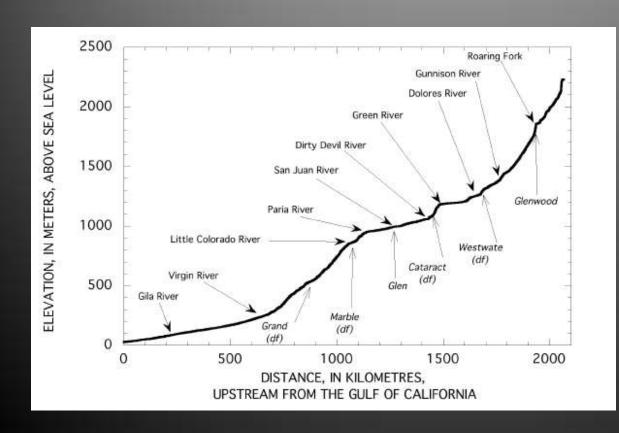
54% of total runoff at Lees Ferry in already in the channels of

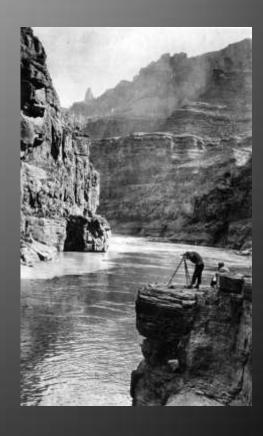
sediment

(data adapted from Iorns et al., 1965)



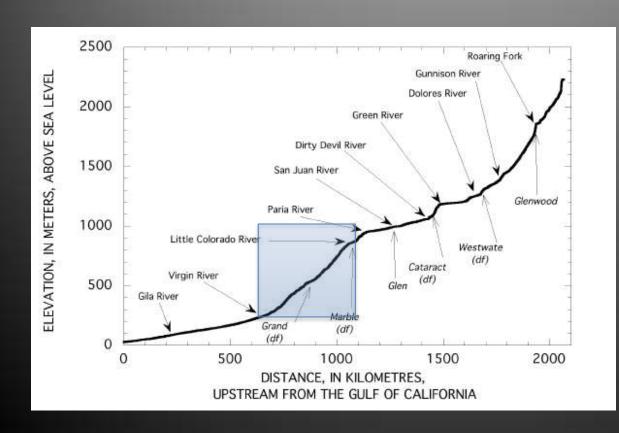
Early measurements of the physical structure of the river date to the 1920s

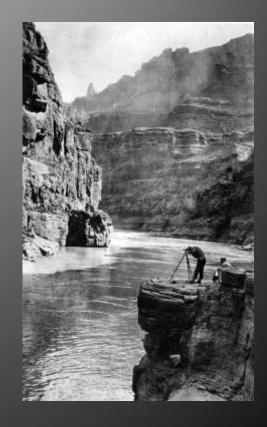


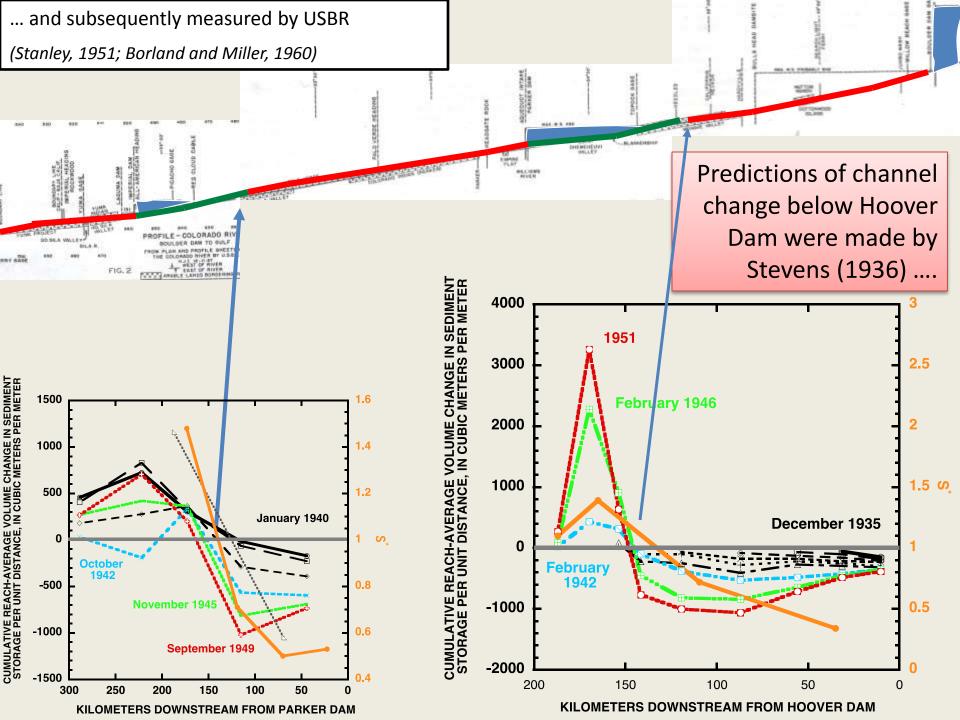




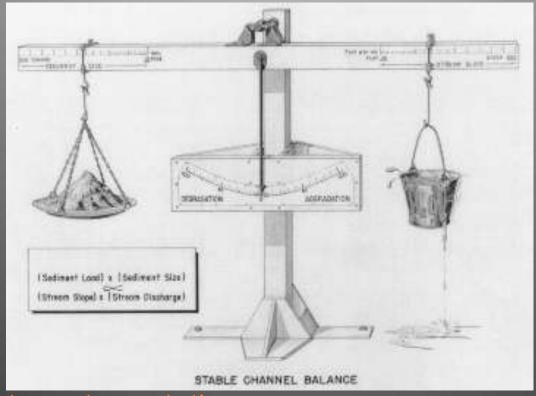
The Grand Canyon segment was the last part of the river system surveyed.







Downstream channel change results from perturbing the balance between the capacity of a river to transport sediment and the amount of sediment supplied to the channel



Factors that induce degradation below dams:

Reduced sediment supply

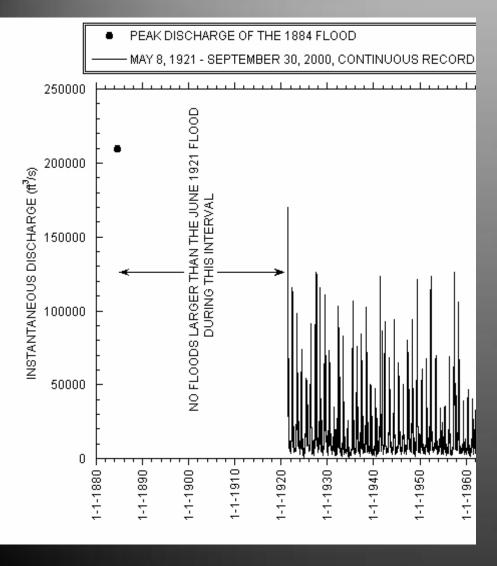
Fining of sediment supply

Increased sediment transport capacity from elevated baseflows

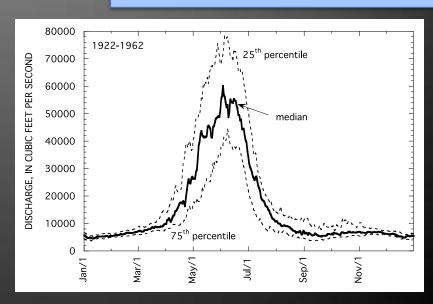
Factors that induce aggradation below dams:

Reduced sediment transport capacity from reduced floods

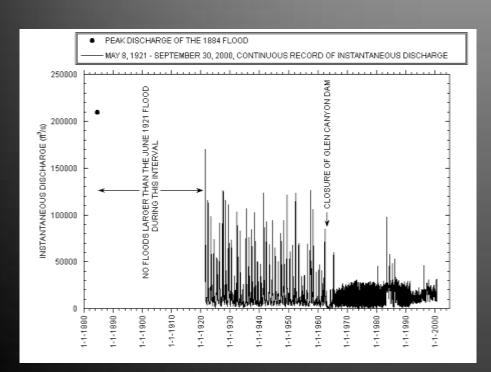
The stream flow that passes through Grand Canyon comes from the Rocky Mountains.

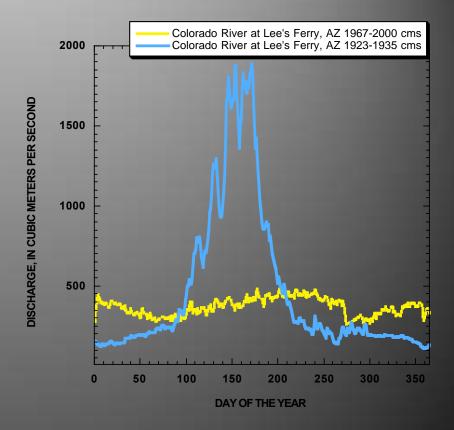


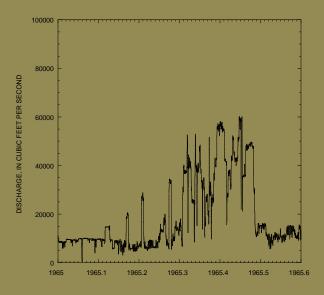
- Pre-dam (1921-1963) flow regime in Grand Canyon
 - **7,980** ft³/s (50% of the time)
 - Flow greatest in June (51,200 ft³/s 50% of the time)
 - Flow lowest in January (5,140 ft³/s 50% of the time)
 - 50,000 ft³/s flood occurred every year, on average
 - 125,000 ft³/s flood occurred every 8 years, on average

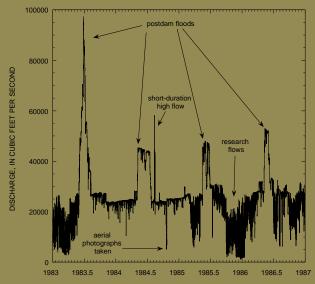


- Post-dam flow regime (1963-2000)
 - 12,600 ft³/s (50% of the time)
 - Flow greatest in August (16,400 ft³/s 50% of the time)
 - Flow lowest in October (10,200 ft³/s 50% of the time)
 - 30,000 ft³/s flood occurs every year, on average
 - 50,000 ft³/s flood occurs every 8 years, on average

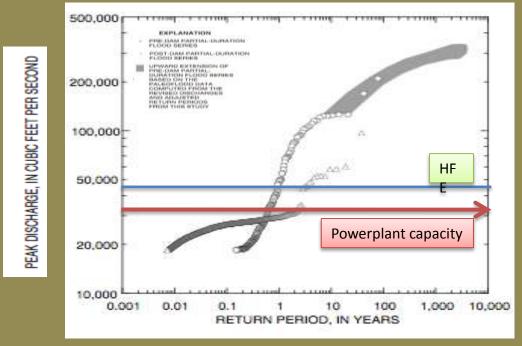






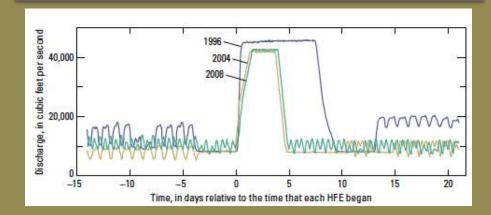


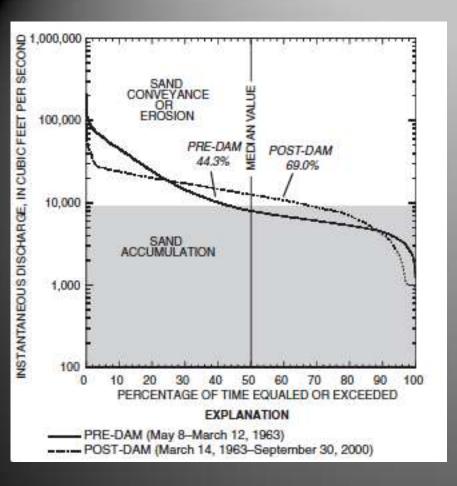
Post-dam floods occurred in 1965, 1983, 1984, 1985, 1986



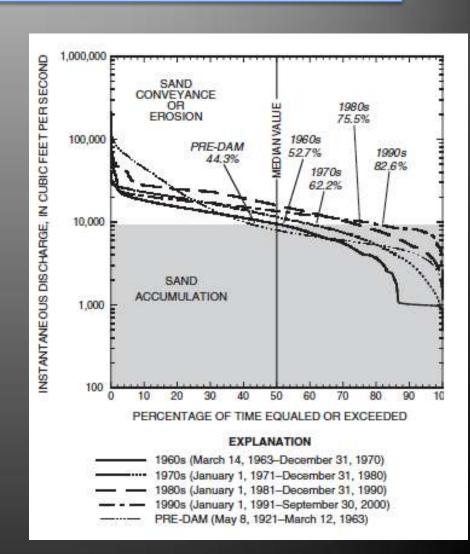
(Topping et al., 2003)

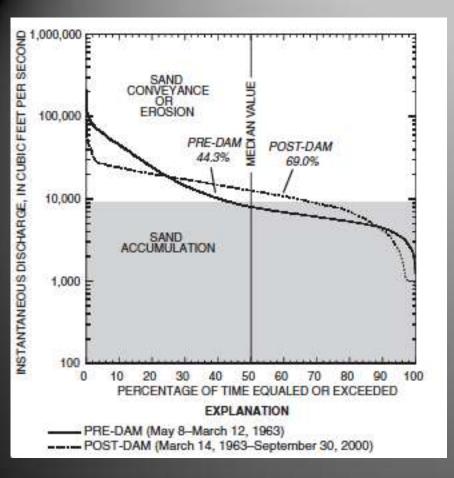
Short duration high flows occurred in 1965, 1980, 1996, 2004, and 2008





Elevated base flows have a significant impact on sand accumulation in Grand Canyon

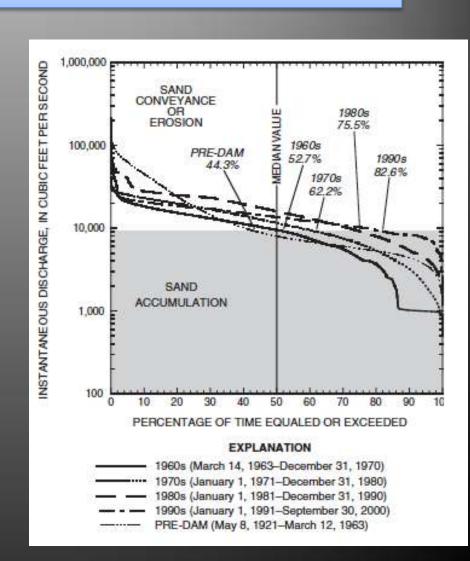


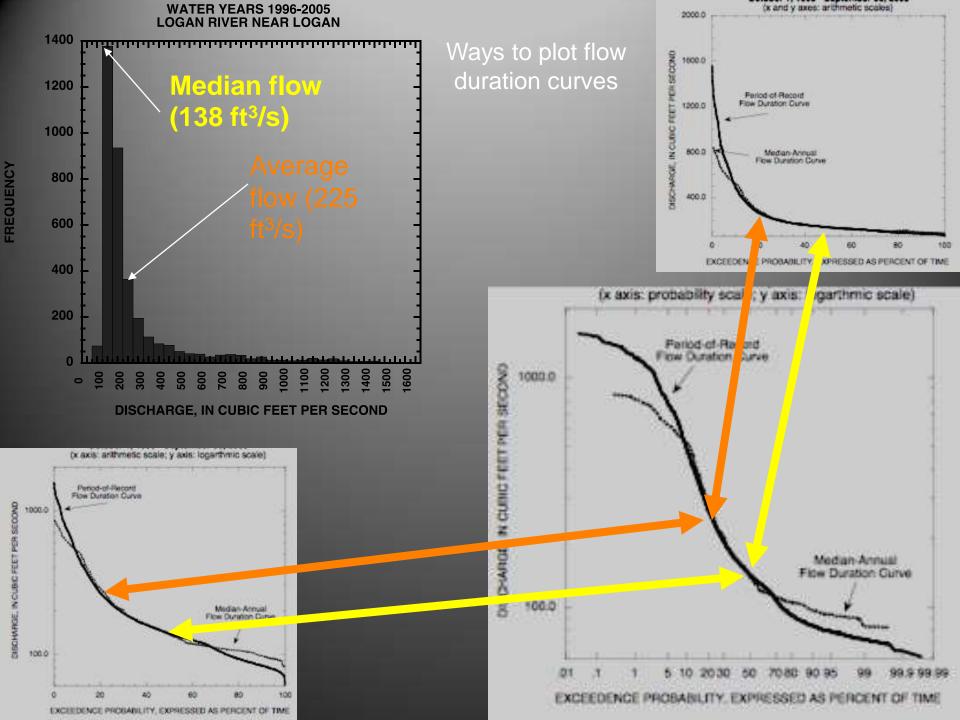


<u>Flow duration curve</u>: a concise picture of the temporal variability of stream flow [but do not reflect seasonality or autocorrelation]

A cumulative frequency curve showing the fraction of time flows exceed a specific value

Elevated base flows have a significant impact on sand accumulation in Grand Canyon







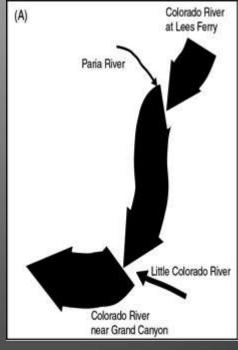
General agreement in the fine sediment supply rate from Paria (1.4-1.9 x 10^6 tons) and Little Colorado River (3.3-3.4 x 10^6 tons)

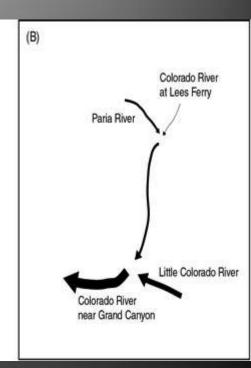
Disagreement about sediment contribution from lesser tributaries:

4.4 x 10⁶ tons (Howard and Dolan, 1981) 0.7 x 10⁶ tons (Randle and Pemberton, 1987)

It is essential to understand the natural sediment supply

Annual average sediment load downstream from Glen Canyon Dam





adapted from Topping (2000)

Longstanding research themes in Grand Canyon research

Is the Colorado River in Grand Canyon evacuating or accumulating sediment?

What controls the large-scale organization of the Colorado River and its valley?

What is the small scale organization of the river?

What are the hydraulics of flow in fan-eddy complexes and how do they change with discharge? How do these changes affect resources?

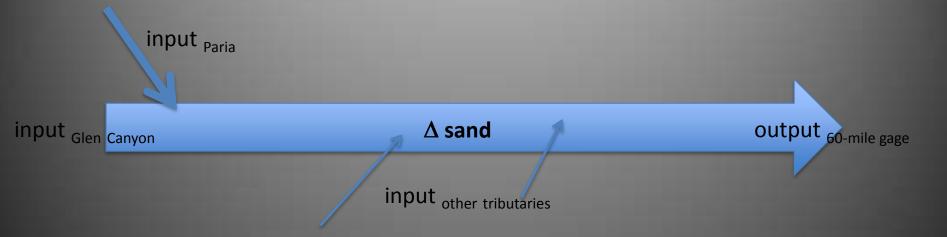
How does the channel bed adjust at annual and decadal timescales?

How do eddy sandbars adjust at annual and decadal timescales?

What have been the short and long-term changes in eddy sandbars?

What is a sediment mass balance?

Inputs – outputs = change in storage



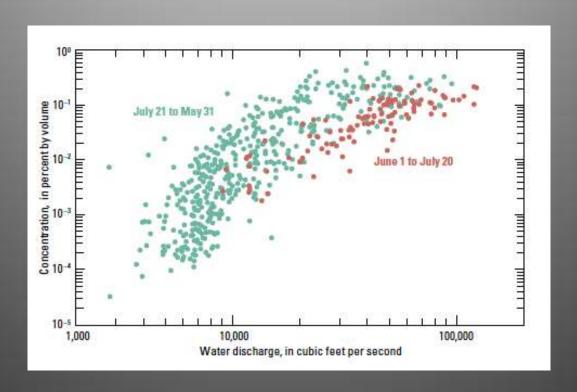
A sediment mass balance for Marble Canyon ...

input $_{\text{Glen Canyon}}$ + input $_{\text{Paria}}$ + input $_{\text{other tributaries}}$ – output $_{\text{60-mile gage}}$ = Δ sand $_{\text{bed}}$ + Δ sand $_{\text{eddies}}$ + Δ sand $_{\text{channel margins}}$

Inputs and outputs = f (water discharge, characteristics of the sediment available to be transported)

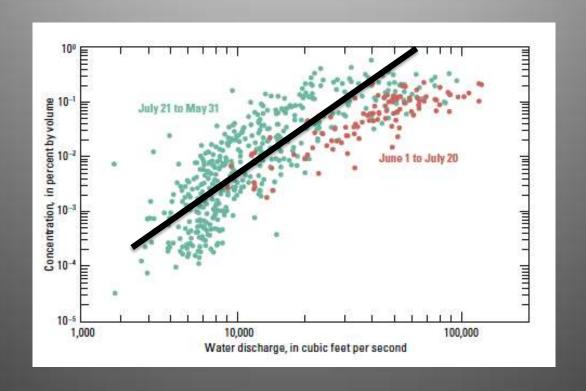
 Δ sand is based on measurements of the places where sand collects

Understanding import and export of fine sediment

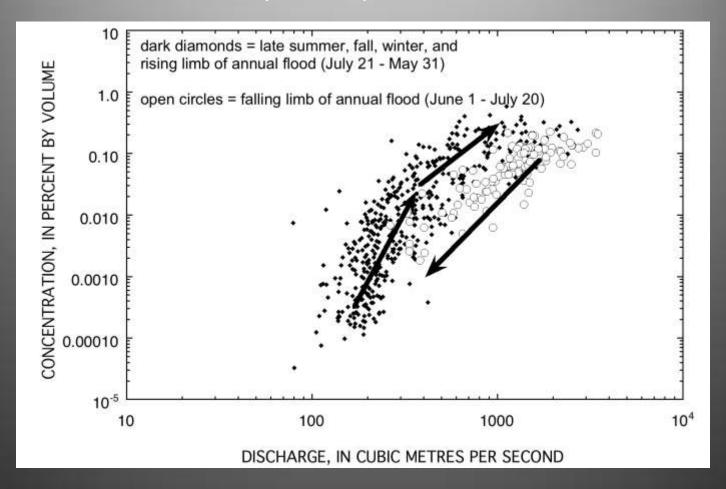


The relation between discharge and sediment transport has wide scatter

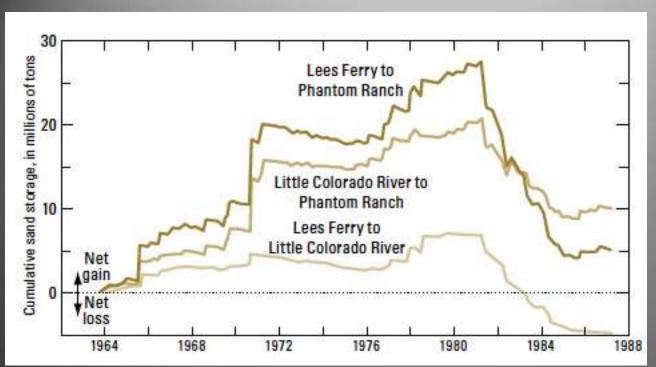
One strategy is to assume the scatter in the relation is random and fit one relation to the data



But there is a systematic pattern to these data ...



River scientists have debated the significance of these patterns for > 50 years *Colby (1964)*

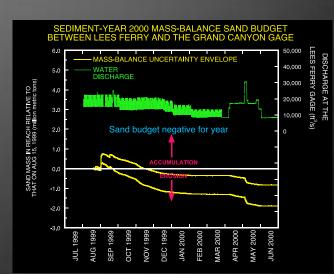


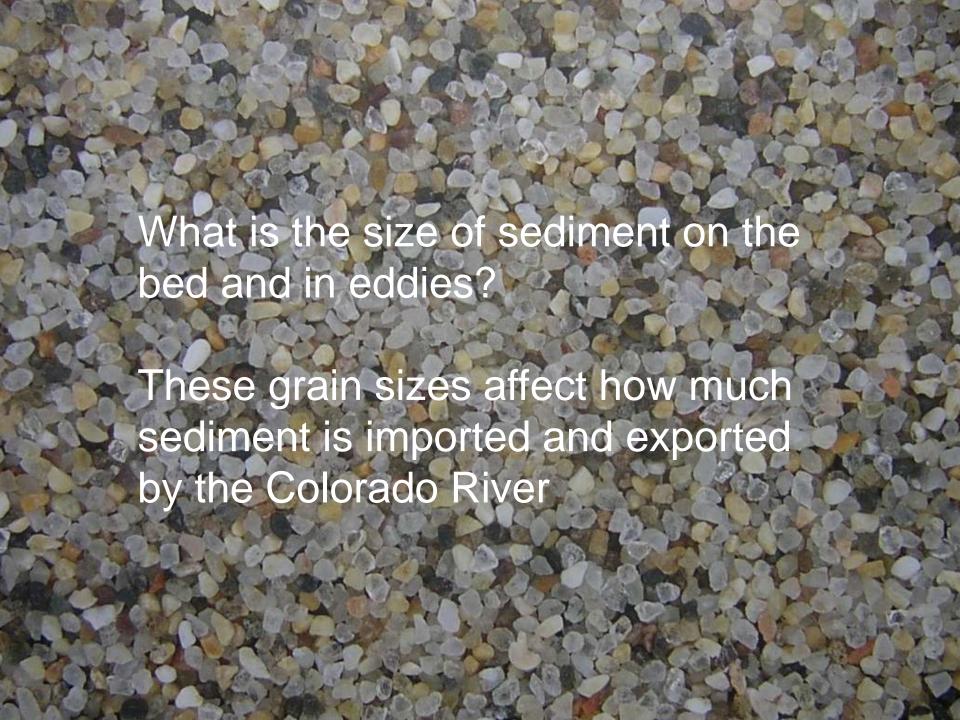
The old view of sediment mass balance

modern view

Failure to account for hysteresis in sediment transport relations led to the wrong conclusion about the sediment mass balance to the post-dam river.

The 1995 EIS argued that the post-dam river had been accumulating sediment when we now know that the river was losing sediment.







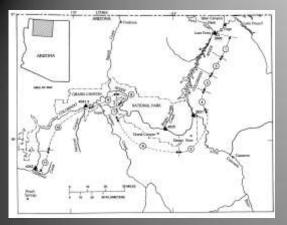
New technologies have been developed

Camera with macro lens and light ring in waterproof housing.

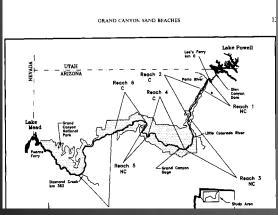
Rubin et al., 2007, Sedimentary Geology.







Schmidt and Graf, 1990



Kearsley et al., 1994

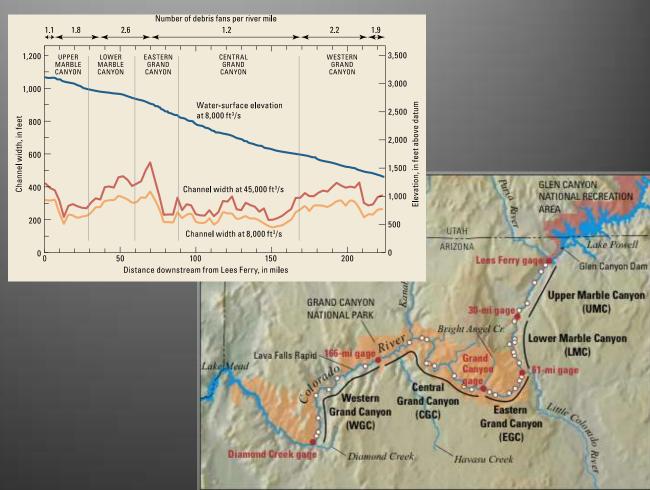
piscak 1980, Stephens and Shoemaker 1987, Schmidt and Graf 1990, Webb et al. 1991).

Development of water supplies has led to regulation of many of the world's rivers (Petts 1984, Gore and Petts 1989), and large dams are typically constructed in bedrock canyons. Impoundments destroy upstream riparian habitats (e.g., Woodbury 1959, Ohmart et al. 1988) and often reduce differences between baseflow and flood stage, increase daily flow fluctuations, reduce sediment transport, and alter existing downstream riparian vegetation composition (Baxter 1977, Turner and Karpiscak 1980, Howard and Dolan 1981, Nilsson 1984, Petts 1984, Williams and Wolman 1984, Ohmart et al. 1988, Johnson 1993). Hourly varying discharges produced by hydroelectric power generation create daily "tidal" fluctuations that are accentuated in narrow

ed rivers i ment there gressive m this paper fluvial ma compositic 1991 in re We discus ecology in rivers. Ou body of ser morphic depled with of the Colo Park and Verystem ide

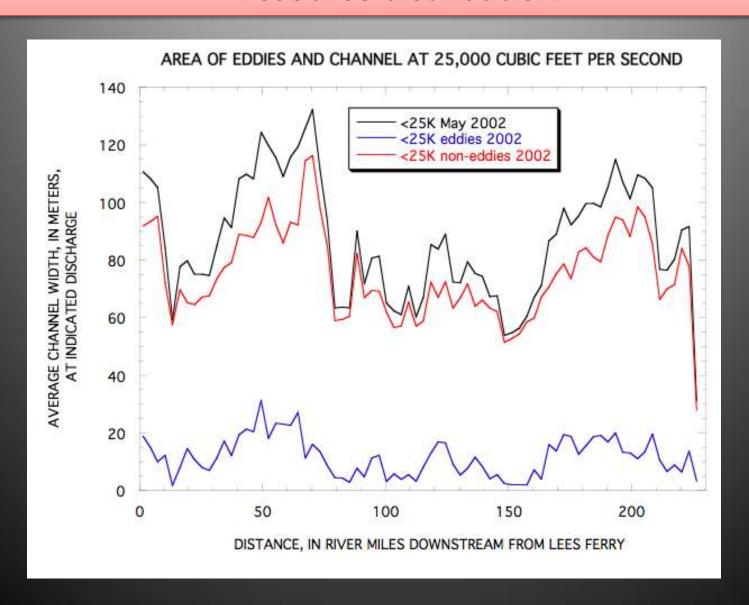
What is the large-scale architecture of the river and its valley?

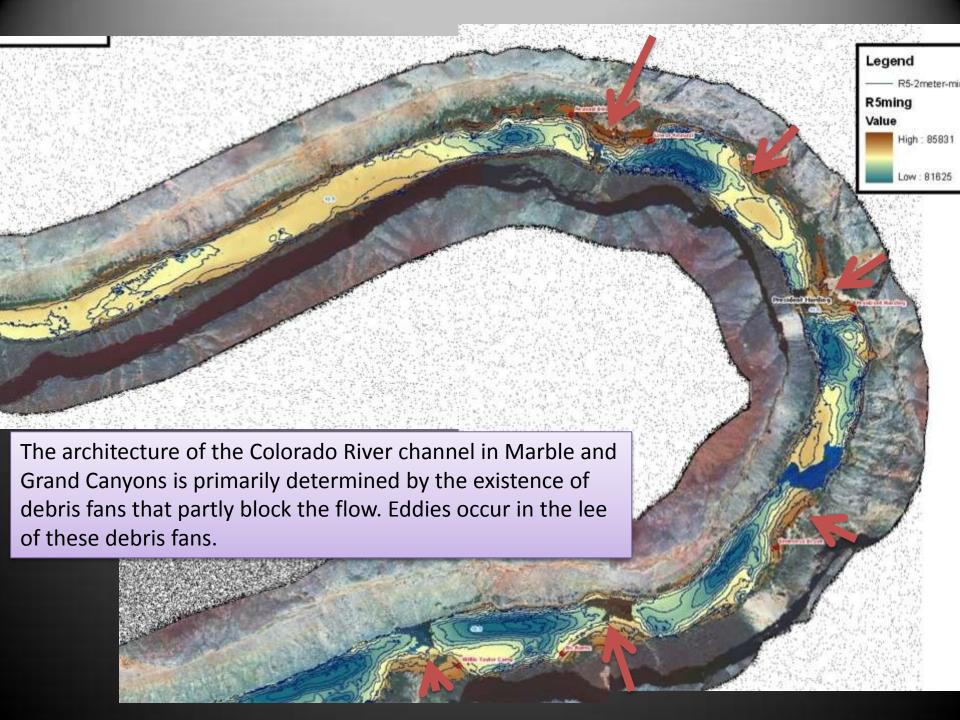
How does that architecture control recreation and ecological values?

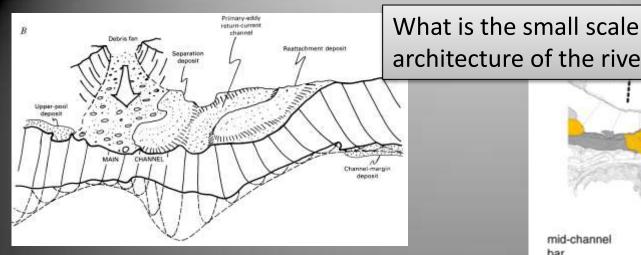


Stevens et al., 1995 Schmidt and Grams, 2011

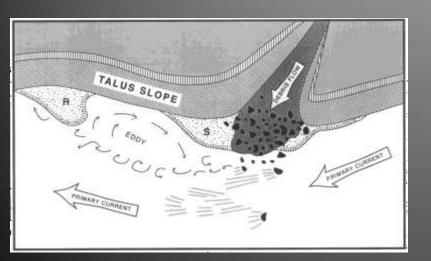
Valley architecture affects large-scale patterns in resource distribution





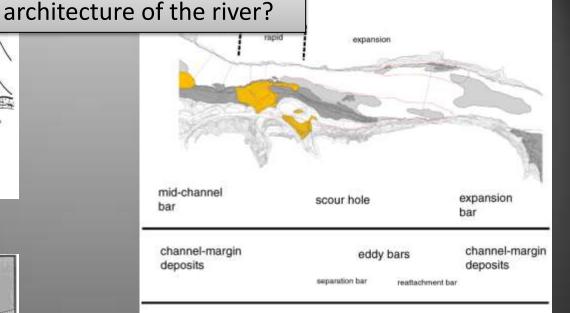


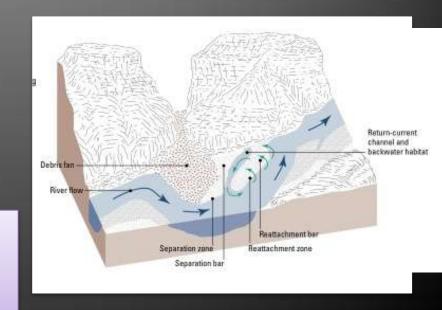
Schmidt and Graf, 1990



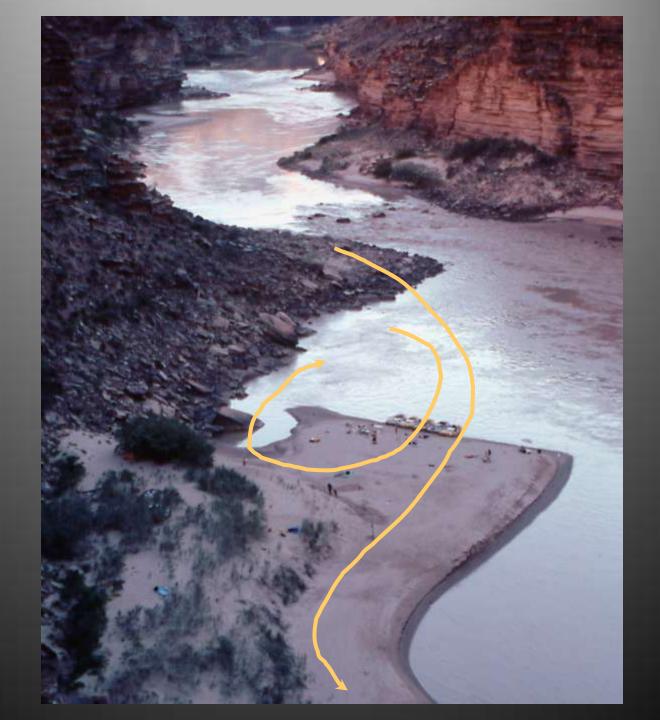
Bauer and Schmidt, 1993

A fan-eddy complex includes (1) an area of ponded flow upstream from a debris fan, (2) a rapid opposite the debris fan, (3) an area where flow width expands and an eddy occurs, and (4) a gravel bar further downstream.

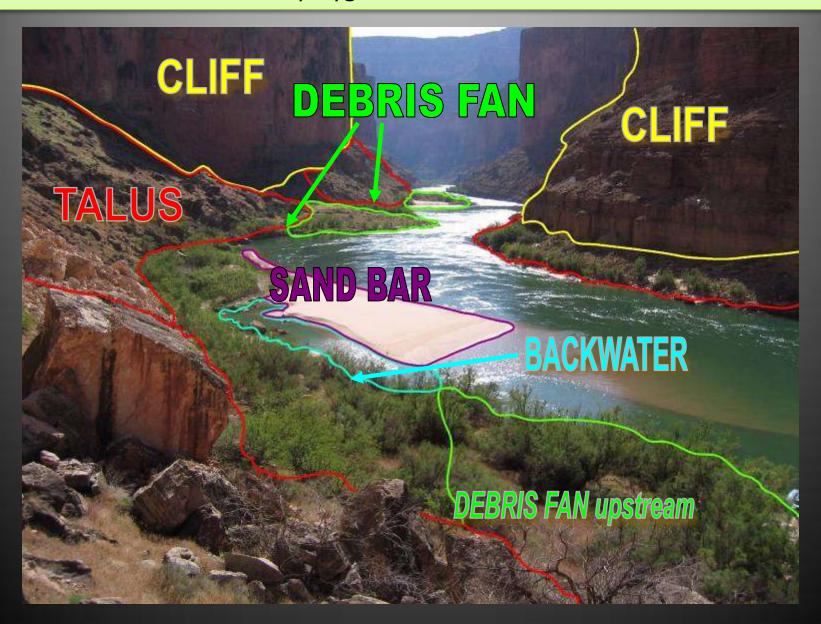


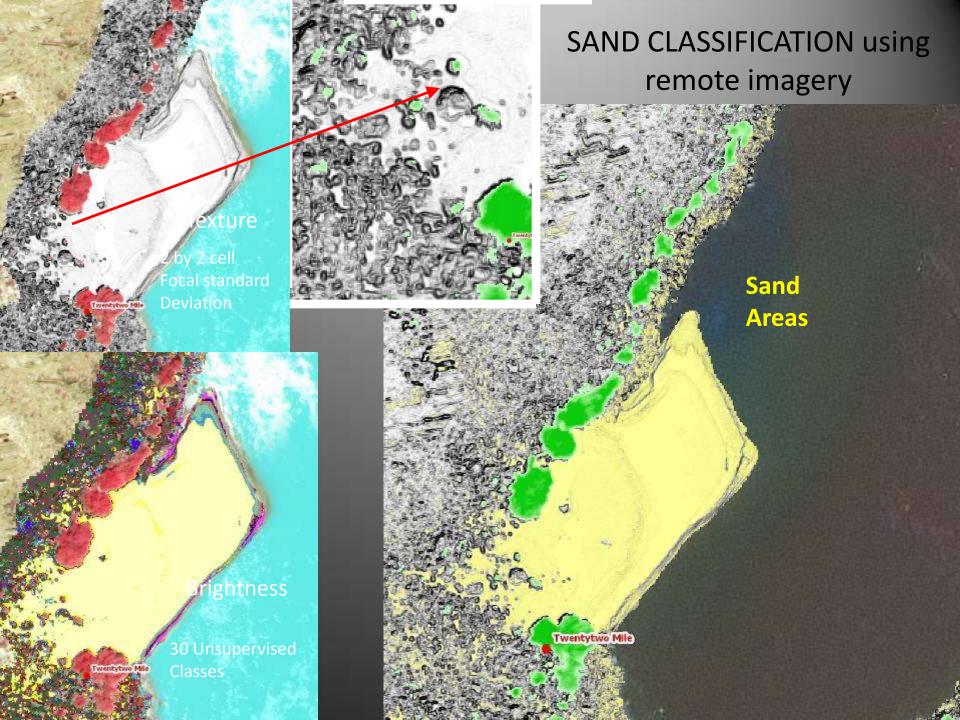


Schmidt and Rubin, 1995



Large-scale geomorphic features can be recognized in aerial photographs and delineated as polygons for the entire river corridor.





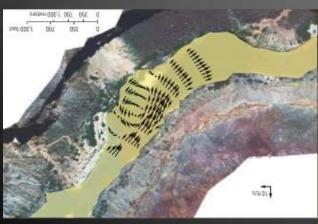
Schmidt and Graf, 1990



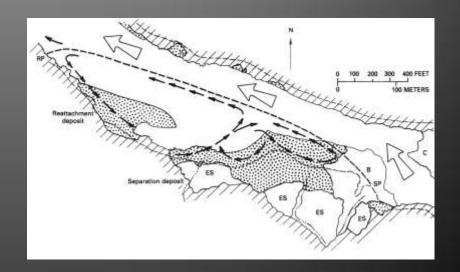
Flow patterns: what are they and how do they change?



Measured depth averaged horizontal velocities at peak flow during 2008 HFE

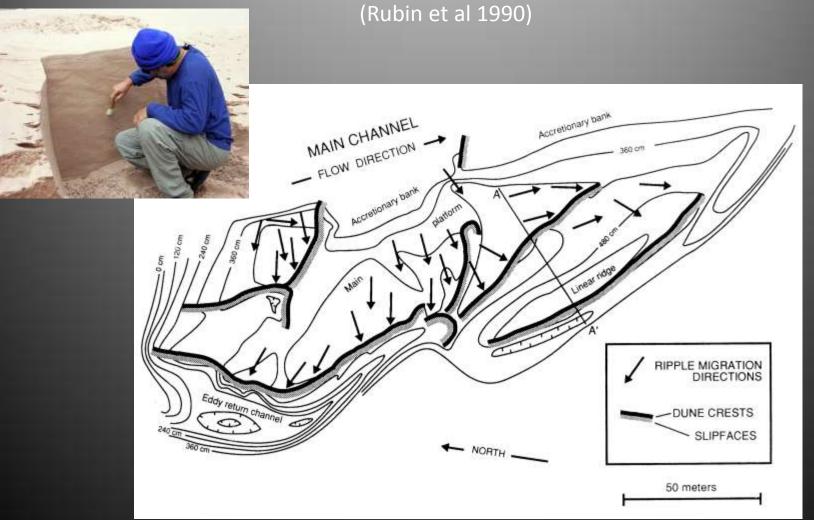


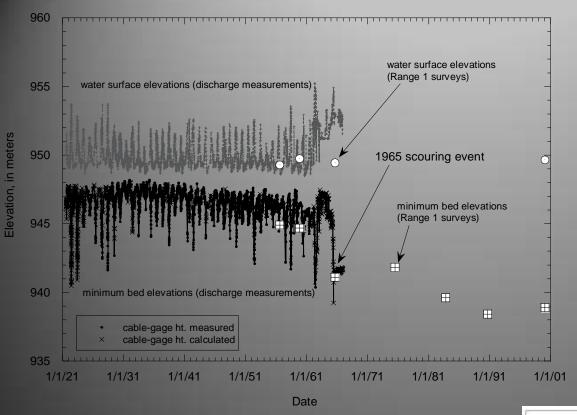
(Wright and Kaplinski, 2010)



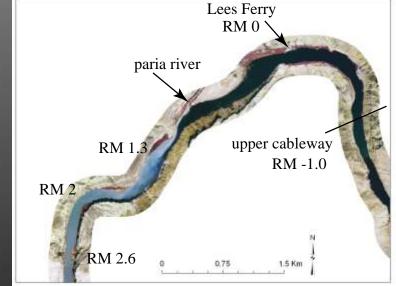
Bar scale features include bar platform and return-current channel

Sedimentary structures reflect flow field





What have been the longterm changes in the bed of the channel?



Changes at Eminence Break campsite, 57 km downstream from the dam

1935

1952, Kent Frost



Many sand bars have dramatically decreased in size

















At some sites, sand bar changes have been minimal (Grapevine Camp river mile 81.8L)



Late afternoon, August 7, 1976 (~daily mean 9,000 ft³/s)



1300 August 7, 1985 (~21,300 ft³/s)



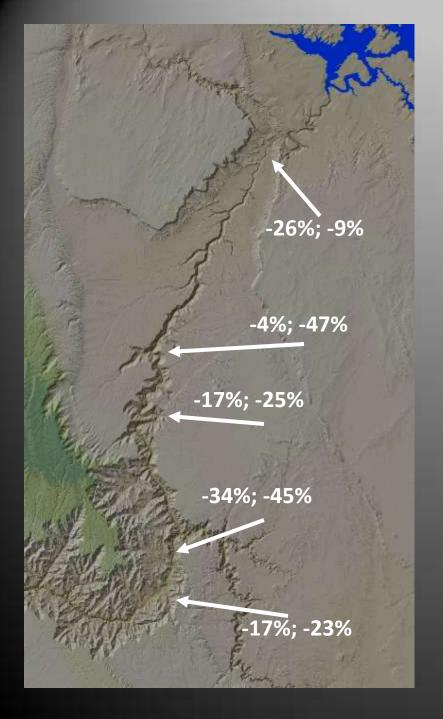
1645 January 24, 1989 (~13,600 ft³/s)



0945 April 6, 2008 (~10,400 ft³/s)

Grapevine, RM 81.76L, Downstream View: 1976, 1985, 1989 sand levels shown in 2008 photo



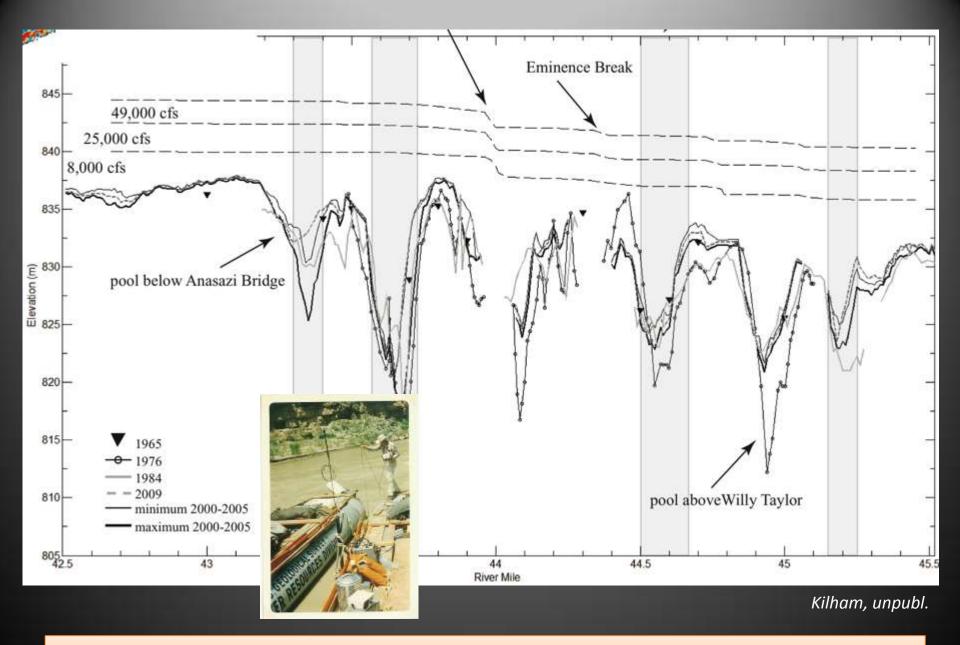


Based on aerial photograph comparison, the average area of eddy bars in the 1990s was at least 25% smaller than the average pre-dam conditions.

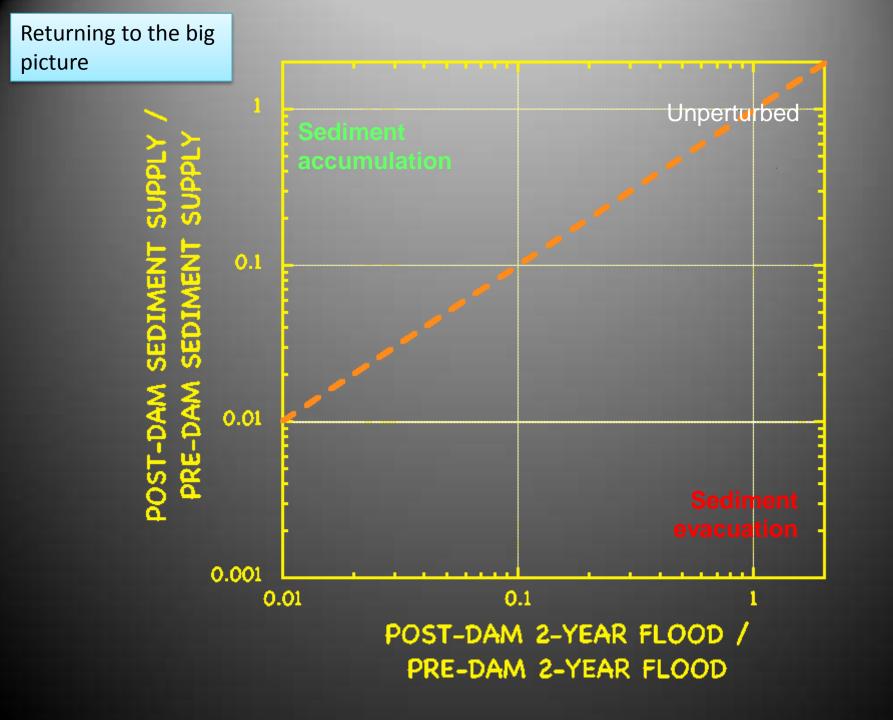
>95% reduction in sand supply has led to ~25% reduction in sand bar size ????

change in the area of all sand

change in sand in the post-dam flood zone

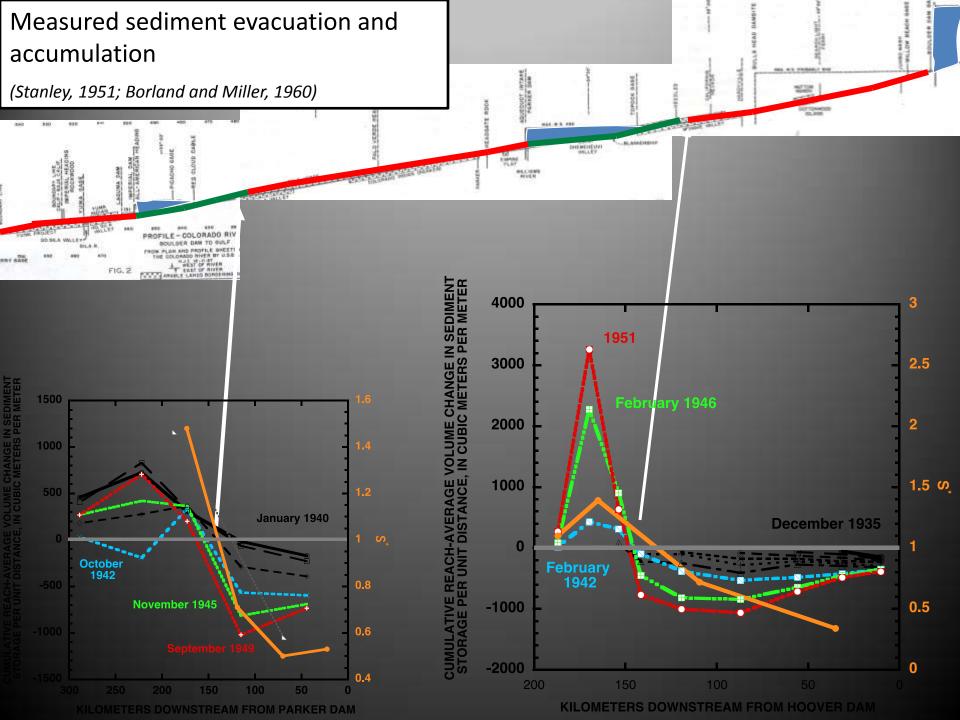


The old longitudinal profiles of the river have been compared with recent traces



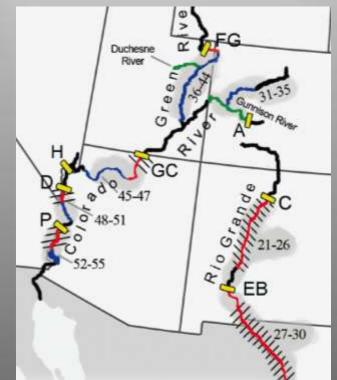
Three Metrics to Describe Channel Change below Dams

- Perturbation of the predam sediment mass balance
 - Assessing shifts towards deficit or surplus
- Post-dam bed incision
- Potential for changes in width

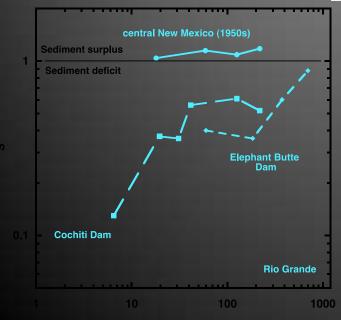


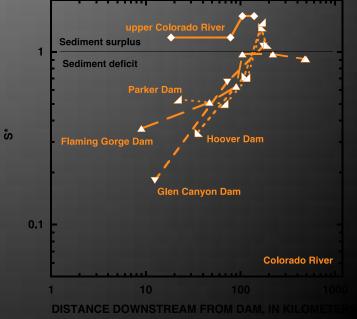
Perturbation of sediment mass balance caused by dams

(hachure indicates zone of bed incision)

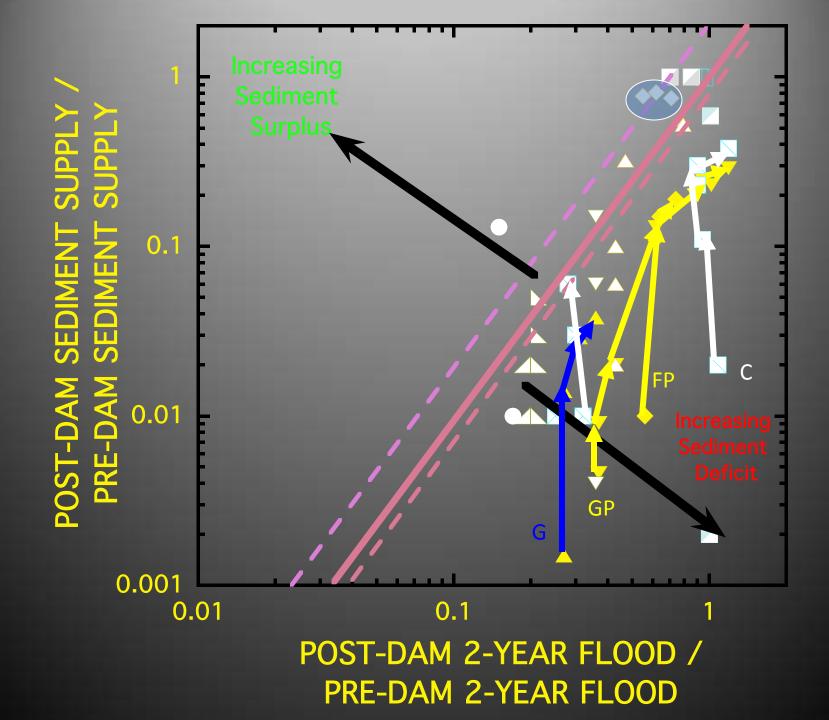


(Schmidt and Wilcock, 2008)





DISTANCE DOWNSTREAM FROM DAM, IN KILOMETERS



Anticipating attributes of channel change

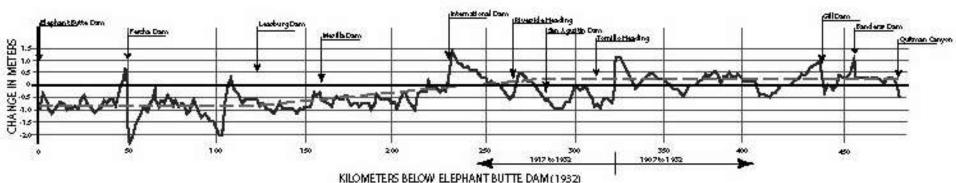
Load Capacity	Load=Capacity	Load>Capacity
Case I	Case 2	Case 3
Case 4 Equal Q	Case 5	Case 6
Case 7	Case 8	Case 9
	Case 1	Case 1 Case 2 Case 4 Case 5



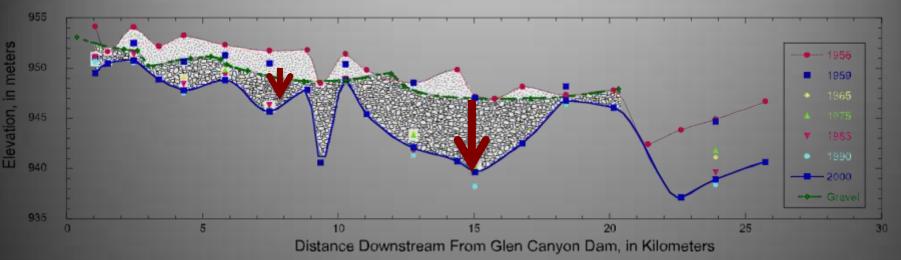
Pattern of channel change downstream from Elephant Butte Dam is typical of the pattern of change below all dams

Downstream from Elephant Butte Dam, the upper river incised its bed as much as 1 m within 225 km downstream from Elephant Butte Dam between ~1917 and ~1933. Downstream from El Paso/Juarez, the bed aggraded about 0.25 m.

(Stevens, 1938)



In Glen Canyon, sediment deficit, bed incision occurred

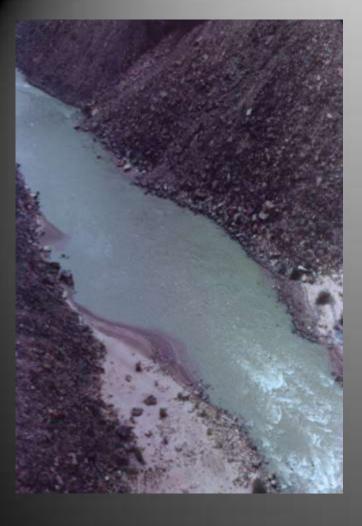


Grams et al., 2007

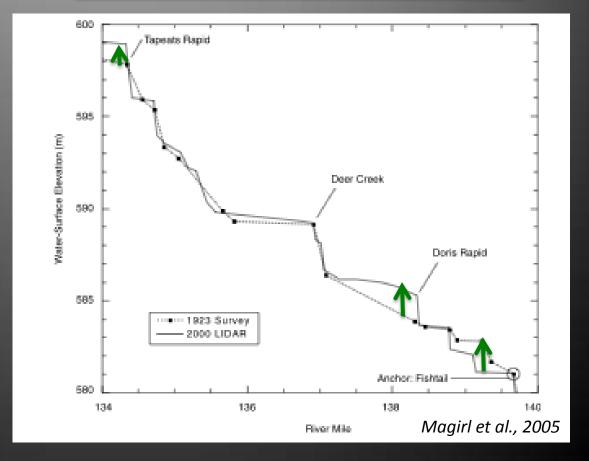
... and conversion of a sand bed to a cobble bed.







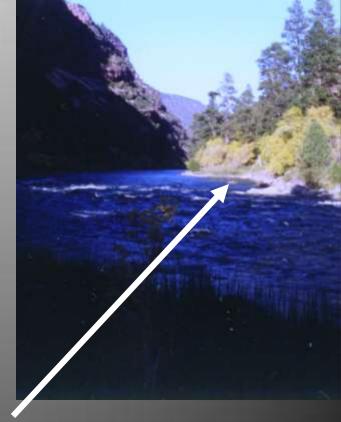
Coarse bouldery rapids prevent bed adjustment in Grand Canyon. With low sediment supply and steep channel slope, mass balance deficit remains and available fine sediment is efficiently removed from system.



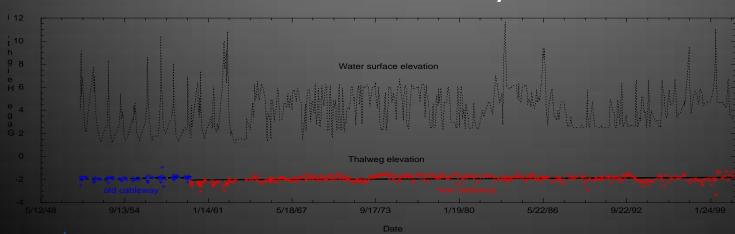
S* = 0.67 (for fine: S* > 1 (for gravel)



Green River immediately below Flaming Gorge Dam



Eroded eddy sand bar



Channels also changes in width

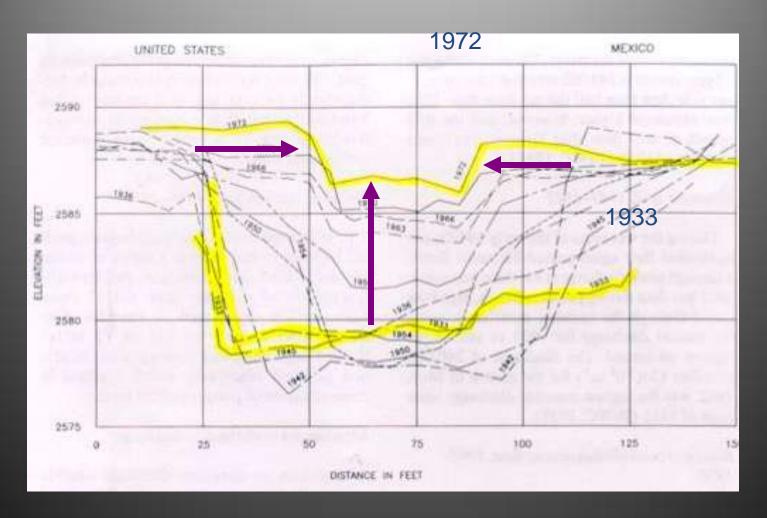
It is difficult to predict the magnitude of narrowing

Narrowing with sediment surplus and no bed incision Narrowing with sediment surplus and bed incision

Narrowing with equilibrium conditions

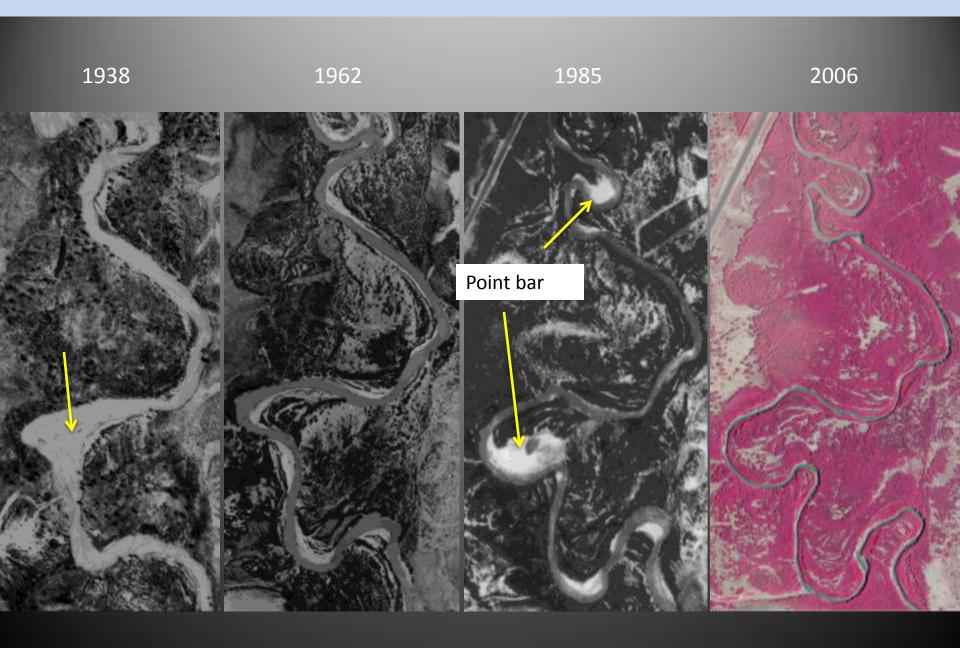
Narrowing with deficit conditions and no bed incision Narrowing with deficit conditions and bed incision

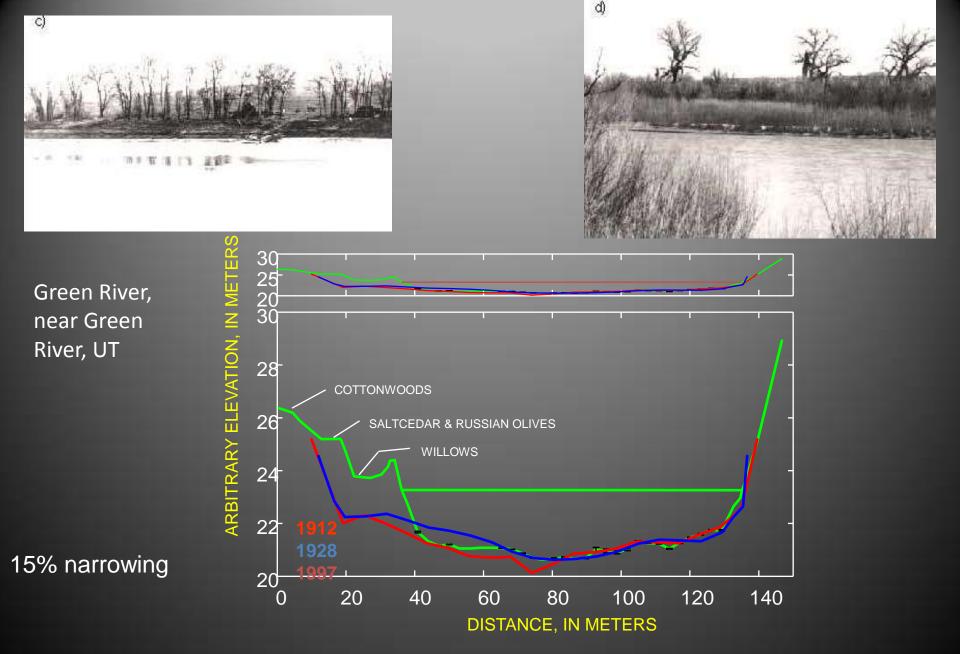
Sediment surplus – narrowing and bed aggradation



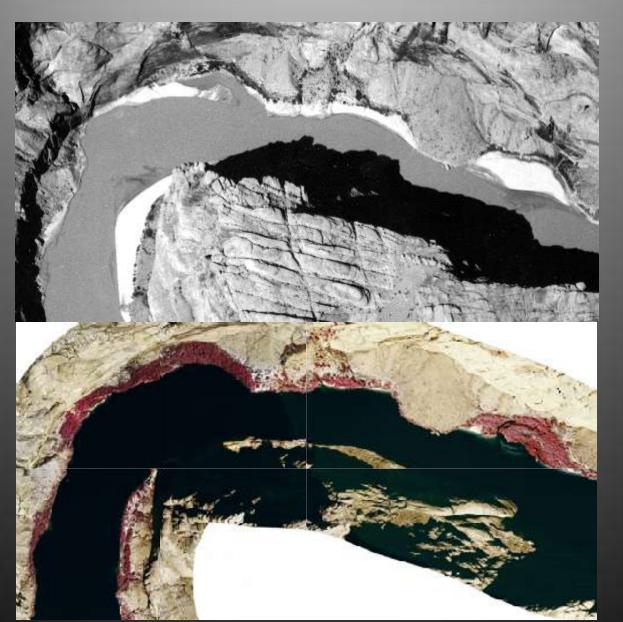
The Rio Grande above the Rio Cochos, near Presidio

Extreme narrowing – San Rafael River $S^* > 1$ with bed incision





These deposits are stabilized by riparian vegetation



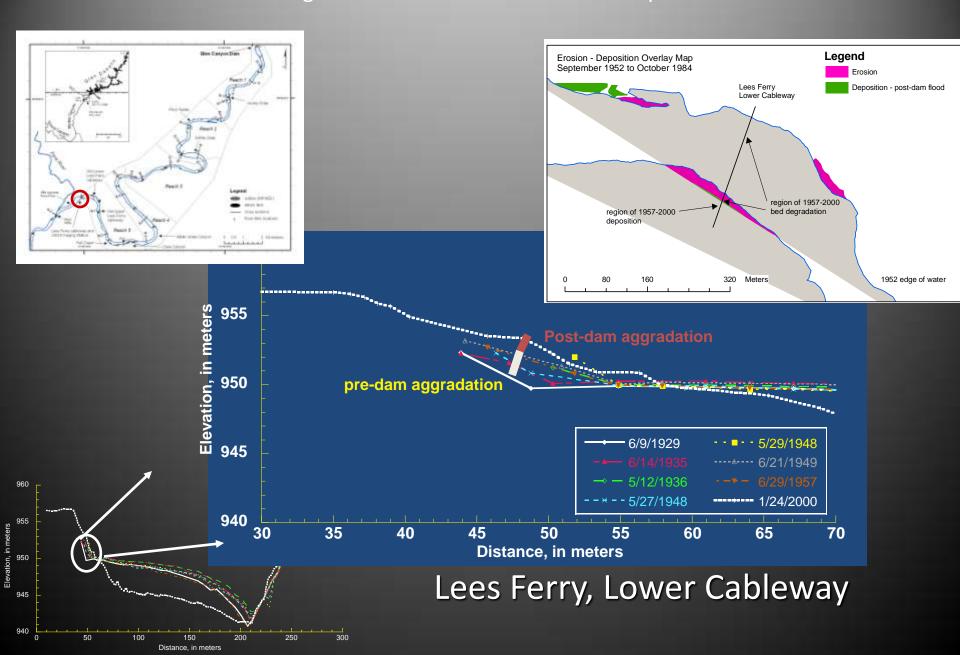
Bed degradation → Perched deposits on channel margins

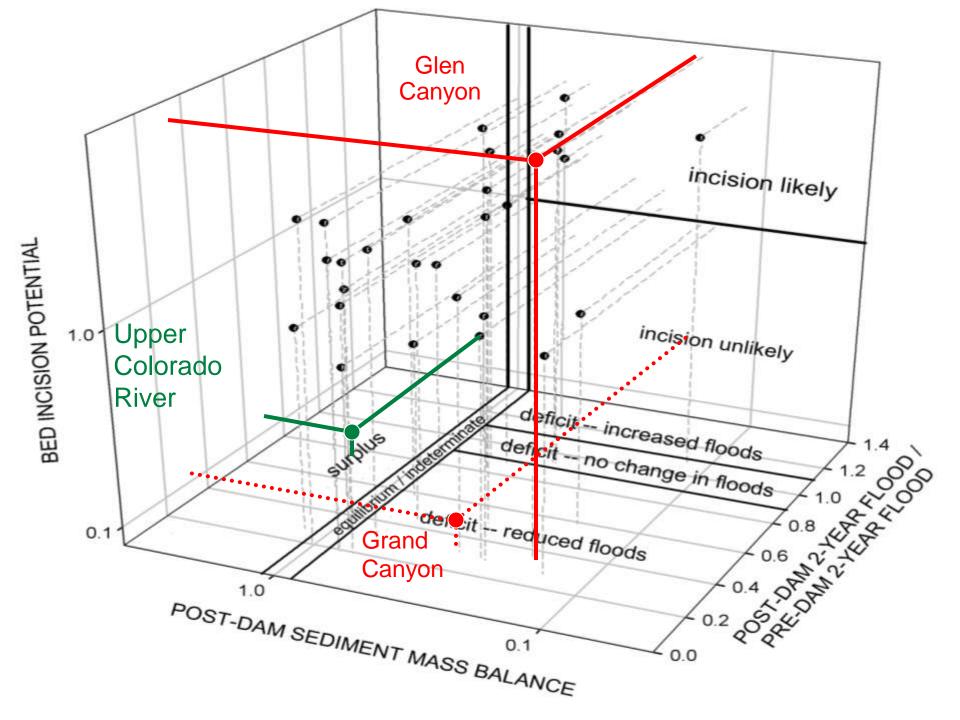


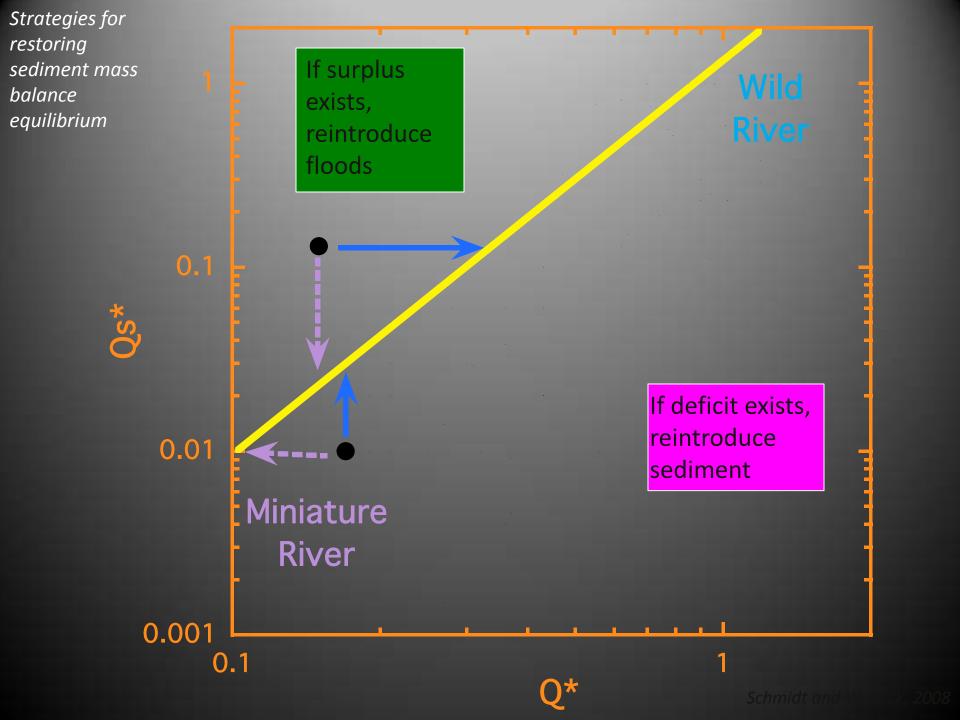


1889

Styles of channel adjustment in segments with degrading bed and **stable** stage-discharge relation: Bed Scour and bank deposition







The costs of adding sediment from Lake Powell into Grand Canyon

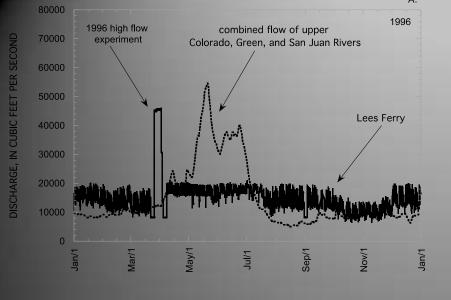
Addition of 4.3 x 10⁶ Mg/yr by dredging and pipeline; *appraisal level cost estimates*

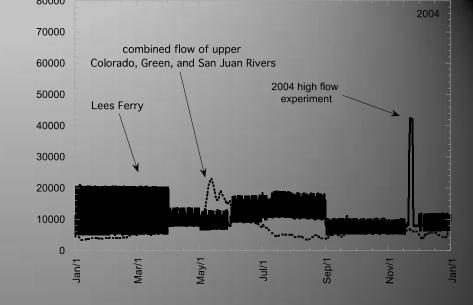
Slurry pipeline Navajo Canyon to Glen Canyon Dam (\$220 million capital costs; \$6.6 million annual operating cost)

Slurry pipeline Navajo Canyon to Lees Ferry (\$430 million capital costs; \$17 million annual operating cost)

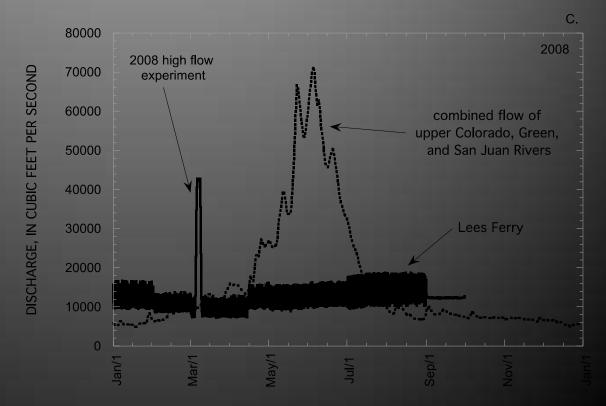


\$44 million/yr is EIS estimate of cost reduced fluctuating flows



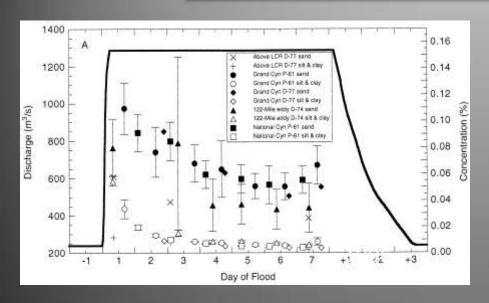


Constraints on reintroducing clear-water floods into a sediment deficit system



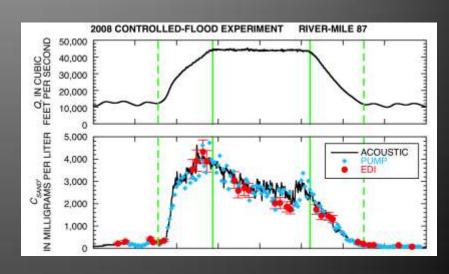
Natural and actual Lees Ferry flows

Controlled floods quickly deplete the available supply.



Change in suspended sediment concentration with time during two large dam releases

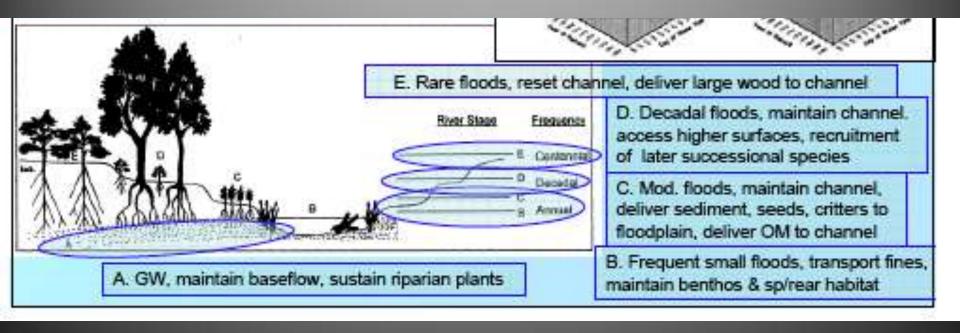
Topping, Rubin, various papers



2008

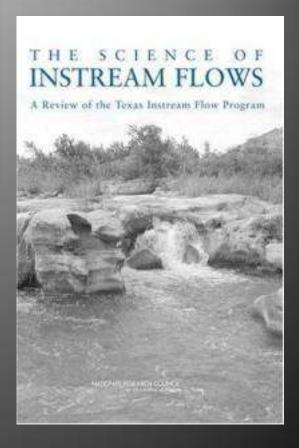
The science of establishing environmental flows

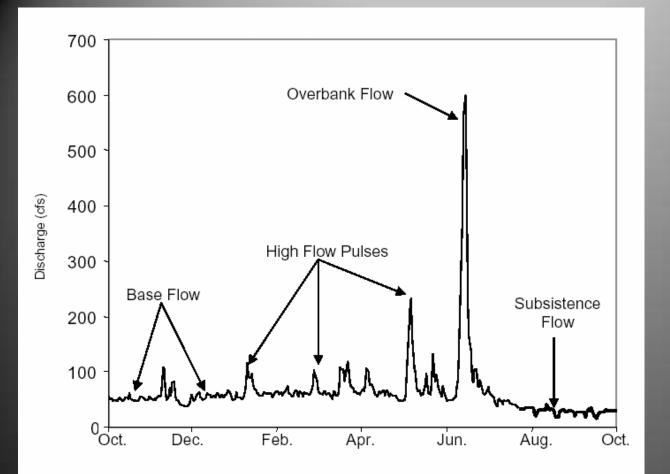
Each part of the flow regime plays a different ecological role: magnitude, duration, frequency, duration, timing and predictability



- Some flows have critical roles in habitat formation and maintenance
 - Overbank flows maintain floodplain features
 - ■Bankfull flows maintain bars, pool/riffle sequences
 - Moderate and high flows transport sediment delivered from upstream
- Flow variability matters because many species evolved to exploit a mosaic of habitats

"... instream flow programs need well-defined and measurable goals to frame instream flow studies and evaluate program progress."





National Research Council, 2005

The Science of Instream Flows: a review of the Texas instream flow program

"state-of-the-science programs use natural flow characteristics as a reference for determining flow needs. Natural river systems have variable flows (also called flow regimes) within a year and among multiple years ... This natural variability is important to sustain aquatic and riparian biota and riverine processes."